

## Transition of SiO<sub>2</sub> in Diatomaceous Earth to Cristobalite

By

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(Received March 20, 1957)

### § 1. Introduction

It is known that silicon dioxide in diatomaceous earth transforms to cristobalite at high temperature.<sup>(1)</sup> One of the present authors (T. N.) studied on the diatomaceous earth from Suzu-shi in the Noto Peninsula, by means of the measurement of electric conductivity.<sup>(2)</sup> In the present report, some theoretical considerations on the electric conductivity data are made, together with another confirmation of this transition by an observation of thermal expansion.

### § 2. Relation between electric conductivity and porosity

On the relation between the electric conductivity of porous materials and their porosity, many experimental investigations have been reported.<sup>(3)</sup> But the results are more or less diverse and the theoretical treatment is difficult.

R. Landauer<sup>(4)</sup> presented an equation giving the electric conductivity of binary metallic mixture  $\sigma_m$  as a function of the conductivities of the components  $\sigma_1$ ,  $\sigma_2$ , and their volume fractions  $x_1$ ,  $x_2$ :

$$\sigma_m = \frac{1}{4} \left[ (3x_1 - 1)\sigma_1 + (3x_2 - 1)\sigma_2 + \sqrt{\{(3x_1 - 1)\sigma_1 + (3x_2 - 1)\sigma_2\}^2 + 8\sigma_1\sigma_2} \right] \quad (1)$$

To apply this equation to porous materials, the variables are transferred as follows:

$\sigma_m \rightarrow \sigma_a$  (apparent conductivity of porous material),

$\sigma_1 \rightarrow \sigma_t$  (true conductivity of the material),

$\sigma_2 \rightarrow 0$  (conductivity of air),

$x_2 (= 1 - x_1) \rightarrow P$  (porosity).

Then a very simple equation is obtained:

$$\sigma_a = \left( 1 - \frac{3}{2} P \right) \sigma_t \quad (2)$$

This equation between  $\sigma_a$  and  $P$  is a straight line cutting  $P$ -axis at  $P=2/3$  and  $\sigma_a$ -axis at  $\sigma_a = \sigma_t$ .

Figure 1 shows the electric conductivity of the diatomaceous earth fired at 900~1200°C for three hours as a function of its porosity, details of measurement being described in reference (2). The solid curve in figure 1 runs along the dotted straight line drawn from a point  $P=2/3=66.7\%$  in the region of  $P=50\sim 60\%$ , which shows that the variation of conductivity with porosity in this range is merely due to the sintering of the sample. For  $P < 50\%$ , the experimental curve departs far from the theoretical one, suggesting that the

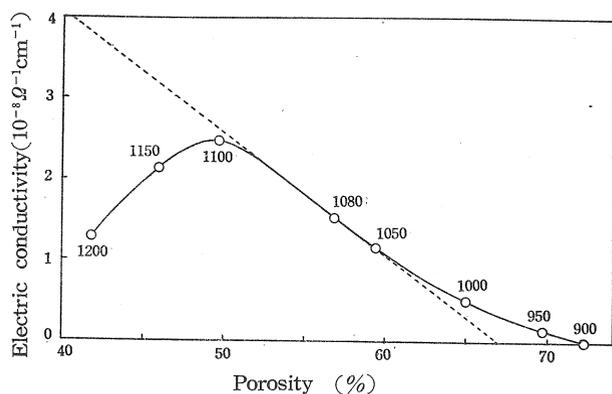


Fig. 1. Electric conductivity at 250°C of fired diatomaceous earth versus its porosity. Firing temperatures are written on the circles in °C. Dotted line shows equation (2).

conductivity in equation (2) vanishes at  $P=2/3$  means that in the porous material, in general, the material particles or crystallites, assumed to be spherical, are scattered separately in the isolated states for the porosity higher than this value, and can no longer have the

change of structure of the material other than sintering is arising. This consideration insures the conclusion of the previous paper.

For  $P>60\%$  also, the experimental curve deviates from the theoretical one. This is interpreted as follows. The equation (1) is based on the assumption that each material particle or crystallite is spherical and can be considered as if surrounded by a homogeneous medium of the mixture. The fact that the

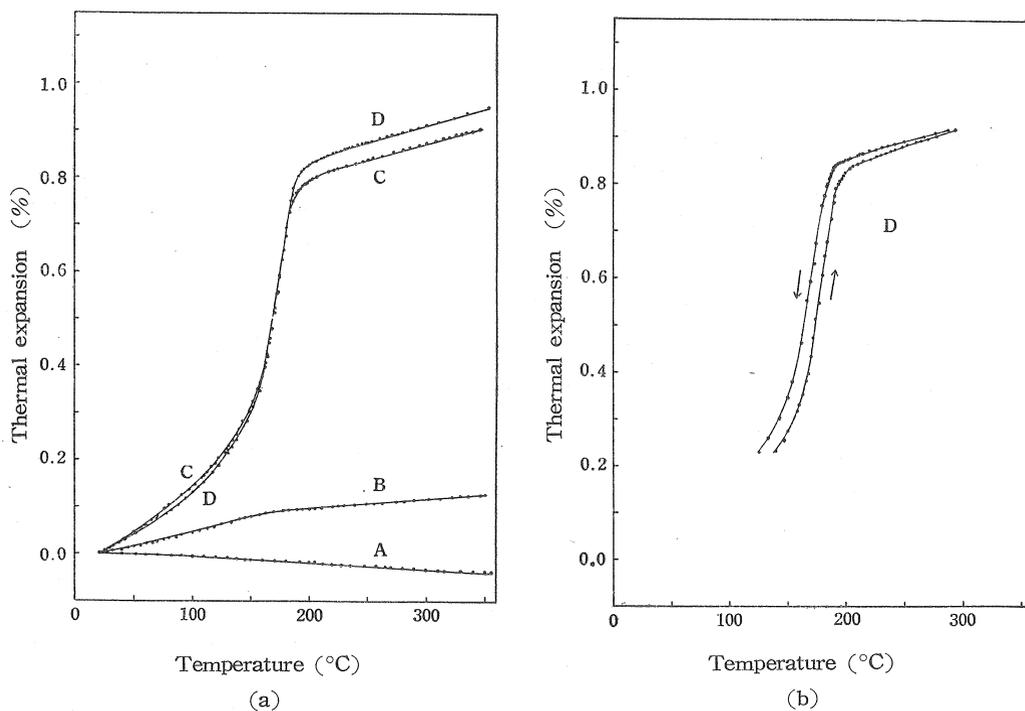


Fig. 2. Thermal expansion of fired diatomaceous earth relative to standard rod.

| Curve | Firing temperature | Standard rod                      |
|-------|--------------------|-----------------------------------|
| A     | 600°C              | Quartz                            |
| B     | 1000 "             | Diatomaceous earth fired at 600°C |
| C     | 1100 "             | " "                               |
| D     | 1200 "             | " "                               |

continuous structure for conducting electricity. In the case of the diatomaceous earth, however, it has so fine reticular structure that it still can be thought continuous even beyond this point, and the conductivity becomes larger than that expected from equation (2). In a word, equation (2) does not hold in such a region of high porosity.

### §3. Observation of thermal expansion

It is well known that cristobalite can exist as  $\alpha$ -cristobalite at room temperature and, when heated, transforms to  $\beta$ -cristobalite at  $180\sim 270^\circ\text{C}$ . Therefore, if the diatomaceous earth fired at high temperature contains much cristobalite, its coefficient of thermal expansion must change at that temperature. This was observed by C. Kawashima.<sup>(5)</sup> The present authors investigated this phenomenon by means of the Honda-Satô's differential dilatometer.

The sample rods of 5.0 or 8.0 cm long and 4.8~5.0 mm in diameter were sawed out from a natural rock. They were fired at 600, 1000, 1100, and  $1200^\circ\text{C}$  for three hours in an electric furnace, then cooled to room temperature, and mounted in the dilatometer. As the standard rod, a quartz rod was used for the sample fired at  $600^\circ\text{C}$ , and  $600^\circ\text{C}$ -fired sample rod, for the samples fired at  $1000\sim 1200^\circ\text{C}$ . The heating rate was about  $1^\circ\text{C}/\text{min}$  in the vicinity of the transition point.

The results are shown in figures 2 and 3. In figure 2, ordinates denote the expansion of samples relative to standard rod divided by their initial length. The gradient of the curve gives the coefficient of differential expansion. In figure 3, the coefficients thus obtained from figure 2(a) are shown for the samples fired at  $1000^\circ\text{C}$  and  $1200^\circ\text{C}$ .

For the sample fired at  $600^\circ\text{C}$ , the thermal expansion is almost linear, or the coefficient is nearly constant, suggesting no transformation to be arising. The sample fired at  $1000^\circ\text{C}$  shows a change of expansion at about  $170^\circ\text{C}$ . It is attributed to the  $\alpha\rightarrow\beta$  transformation of cristobalite slightly contained in the sample. In Kawashima's result the expansion of the sample fired at  $1000^\circ\text{C}$  is almost linear. For the samples fired at  $1100^\circ\text{C}$  and  $1200^\circ\text{C}$ , the change of expansion is distinct, evidently indicating that these samples are abundant in cristobalite. This result is similar to Kawashima's one, but the change is larger for the present samples. The transformation takes place during  $140\sim 190^\circ\text{C}$  for the sample fired at

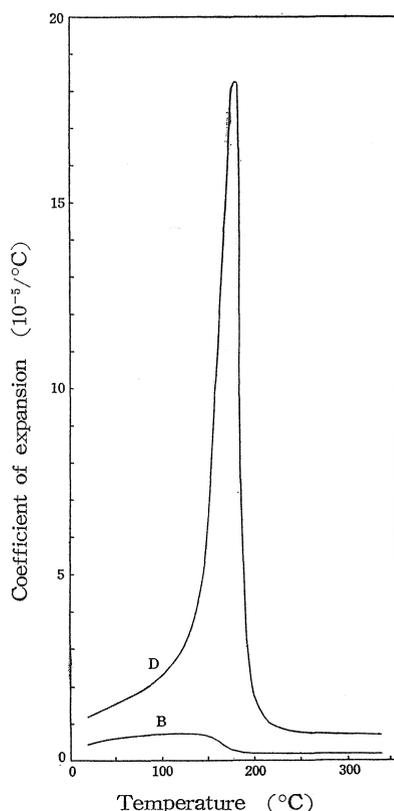


Fig. 3. Coefficient of differential expansion of fired diatomaceous earth estimated from figure 2. Labels on the curves have the same meaning as in figure 2.

1100°C, and during 140~195°C for the one fired at 1200°C. Thus the transition point of cristobalite in diatomaceous earth is somewhat lower than that of pure crystal. It may be an effect due to impurities contained in the diatomaceous earth. The heating curve and cooling curve show a hysteresis in figure 2(b).

#### §4. Summary

(1) The relation between the electric conductivity and the porosity of the diatomaceous earth fired at temperatures below 1100°C satisfies an equation derived from Landauer's one. The samples fired over 1100°C does not satisfy the above relation. This is attributed to the transition of this material to cristobalite.

(2) The thermal expansion of the fired samples was observed by means of the Honda-Satô's differential dilatometer. For the samples fired over 1100°C, the coefficient of expansion changes at 140~190°C corresponding to the  $\alpha \rightarrow \beta$  transformation of cristobalite.

The authors wish to express their gratitude to Prof. H. Shôji for constant guidance in the course of the work.

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